ORIGINAL ARTICLE



Future soil loss in highland Ethiopia under changing climate and land use

Desalew Meseret Moges 1 • Alexander Kmoch 2 • H. Gangadhara Bhat 3 • Evelyn Uuemaa 2

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Abstract

Soil erosion caused by climate and land-use changes is one of the biggest environmental challenges in highland Ethiopia. The aim of this study was to assess the future soil erosion risks and evaluate the potential conservation measures in the Rib watershed, northwestern highland Ethiopia. We used the HadGEM2-ES model with a moderate greenhouse gas (GHG) concentration scenario (RCP4.5) to project the future climate. The future land-use patterns were predicted using the CA-Markov model. We integrated the RUSLE model with GIS to estimate the spatial distribution of soil loss and identify erosion risk areas. We found that the Rib watershed is highly vulnerable to future climate and land-use changes, leading to a high soil erosion risk. Despite slight growth of forest cover during the study period, the total soil loss for the watershed was estimated to be 7.93×10^6 t year⁻¹ in 2017 and was predicted to increase to 9.75×10^6 t year⁻¹ in 2050, an increase of about 23%. The increase in forest cover was due to the expansion of the area of eucalyptus plantations which are more prone to erosion. Moreover, field survey showed that the residual native forests are sparsely vegetated and mostly used for cattle grazing, increasing the erosion risk even more. In contrast, the combined use of afforestation with native trees and physical soil conservation measures in the upper areas of the catchment could decrease soil loss by 62%. Our results stress the importance of combining soil conservation measures, including converting eucalyptus plantations to native forests, to mitigate the effects of future climate change and increased agricultural production on soil erosion.

Keywords Modeling · HadGEM2-ES model · RUSLE model · CA-Markov model · GIS · Nature-based solutions

Introduction

Soil is a vital part of the terrestrial ecosystem that contributes to the delivery of primary ecosystem services (Pereira et al.

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Desalew Meseret Moges mdessalew@gmail.com

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- Department of Geoinformatics, Mangalore University, Mangalagangotri 574199, India
- Department of Geography, University of Tartu, Vanemuise 46, 51003 Tartu, Estonia
- Department of Marine Geology, Mangalore University, Mangalagangotri 574199, India

2018). Healthy soil performs a large number of crucial functions, including the production of food and biomass products, storage and filtration of carbon and water, and regulation of carbon dioxide and other emitted gases (Breure et al. 2018). Therefore, soils must be protected and used in a sustainable way. Keesstra et al. (2018) suggest that nature-based solutions, including solutions that target soils and landscapes, are cost-effective long-term solutions for hydrological risks and land degradation, and can support sustainable soil use.

However, water-induced soil erosion has become a major ecological and environmental problem, as it leads to a continuous reduction of soil quality and productivity in many places around the world (Mekuriaw et al. 2018). Although soil erosion is a global phenomenon, it has become particularly rapid and severe in highland Ethiopia due to complex interactions among anthropogenic and natural factors. Annually, Ethiopia loses more than 1.5×10^9 t of topsoil from the highlands due to erosion, and this soil could have added about 1.5×10^6 t of grain to the country's annual harvest (Tamene and Vlek 2008).

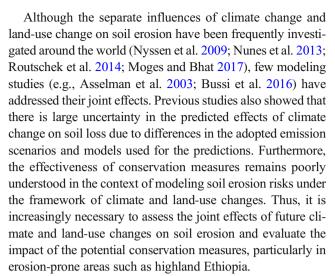


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This makes soil erosion one of the biggest environmental constraints to agricultural sustainability, food security, and rural development in highland Ethiopia (Teshome et al. 2016).

Climate and land-use changes are among the most important agents that are causing soil erosion in highland Ethiopia (Tadesse et al. 2017; Bekele et al. 2018). Deforestation and the expansion of cropland in Ethiopia have already reached the highest level observed in the past several centuries. Currently, cropland in most parts of the Ethiopian highlands occupies more than 80% of the total land, whereas natural forest covers less than 3%, and much of this forest is significantly degraded. About 40% of Ethiopia's total land was once covered by various types of natural forest, but this proportion had shrunk to 16% in the 1950s and continued to decrease, reaching 2.7% in the 1990s (FAO 2010). Cebecauer and Hofierka (2008) and Mohammad and Adam (2010) note that such a rapid reduction in forest cover significantly increases surface runoff and soil erosion in various ways. Forest reduces the rate of soil erosion by stabilizing soils and enhancing their water-holding capacity. Thus, any kind of forest destruction directly increases soil losses by facilitating surface runoff. Land-use change also markedly affects soil organic carbon through its effect on humification, decomposition, and mineralization of soil organic matter, which in turn influences the rate of soil loss because organic matter increases soil cohesion and water retention (Basaran et al. 2007). Therefore, by properly adjusting land-use patterns, soil properties can be greatly improved and runoff and soil erosion rates can be reduced (Zhang et al. 2008).

In addition to land-use change, climate change represents one of the critical global environmental issues that is affecting hydro-meteorological processes and increasing surface runoff and soil erosion. One of the most direct effects of climate change is due to a change in the erosive power of the rainfall (Pruski and Nearing 2002; Simonneaux et al. 2015). Higher rainfall modifies the hydrological conditions in a given area and influences the infiltration and runoff rates (Pruski and Nearing 2002), which in turn causes an increase in surface runoff and soil erosion even when land use has not changed (Klik and Eitzinger 2010). Climate change also determines plant biomass production and decomposition, evapotranspiration, soil microbial activities, and changes in land use (Bussi et al. 2014; Williams et al. 1996), all of which control the rate of soil loss in different ways. Unless proper soil protection practices are carried out, the increasing rate of rainfall erosivity is predicted to cause an extraordinary increase in soil erosion worldwide, with serious implications for economic, social, and environmental sustainability (Plangoen et al. 2013; Paroissien et al. 2015). The economic implications of climate change are particularly severe in developing countries such as Ethiopia for many reasons, including high dependence on rainfed agriculture, low adaptive capacity, and high ecological and social vulnerability (Nigussie et al. 2018).



Several empirical models have been developed to assess the response of soil erosion to changing climate and landuse conditions. The revised universal soil loss equation (RUSLE; Renard et al. 1997) is among the most widely used empirical models due to its structural simplicity (Volk et al. 2010), compatibility with many geospatial technologies (Bhandari et al. 2015), and feasibility for studies at a watershed scale (Prasannakumar et al. 2012). The aim of the present study was to evaluate the future impacts of climate and landuse changes on soil loss in a representative watershed (the Rib watershed, in highland Ethiopia) using the RUSLE model integrated with geographical information system (GIS) software. We also assessed the effectiveness of several conservation measures in reducing soil erosion rates. We selected the Rib watershed as a case study because it is one of the most severely eroded watersheds in highland Ethiopia, mainly due to changes in rainfall erosivity, land-use changes, the widespread use of intensive agriculture, and limited use of conservation measures.

Materials and methods

Study area

This study was conducted in the Rib watershed, which is located in the northwestern highlands of Ethiopia (Fig. 1). It covers an area of about 1975 km², and 376,256 residents rely on agrarian activities for their livelihood. The watershed's population is growing, which will increase the pressure on agricultural land. The watershed's topography is generally rugged and mountainous, with the elevation ranging from 1758 m above sea level (m a.s.l.) in the western part to 4104 m a.s.l. in the southeastern part. The watershed has a climate ranging from humid to sub-humid and contains several ecosystems and resource types. The average annual precipitation from 1973 to 2016 was approximately 1502 mm



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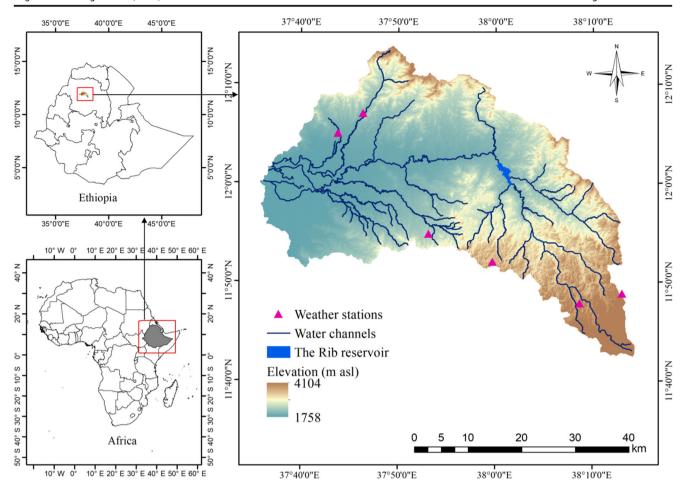


Fig. 1 Location and topography of the study area

(Moges and Bhat 2018). The area receives the maximum amount of rain between June and September (the main rainy season) and the minimum amount from October to February (the dry season), when there is little or no rain. The mean annual temperature ranges from 15 to 17 °C. Moges and Bhat (2018) revealed that about 88% of the watershed is occupied by agricultural land, whereas forest covers only 3.5% of the catchment. The major rainfed cereal crops grown in the area include barley, wheat, tef, maize, and legumes such as beans and peas (Moges and Bhat 2017). The watershed contains a wide range of soils derived from weathered volcanic rock. Based on the FAO-UNESCO soil classification system, the four major soil types (which account for more than 80% of the soils in the watershed) are Luvisols, Vertisols, Leptosols, and Regosols.

Land-use data sources and classification

We analyzed the land-use change patterns and trends in the watershed based on remotely sensed images from three reference years (1986, 2001, and 2017). Time-series Landsat data with a path/row of 169/052 were obtained from the Earth Resources Observation and Science data center (https://

www.usgs.gov/centers/eros) of the US Geological Survey. All images were geometrically co-registered to the Universal Transverse Mercator (UTM) projection coordinate system (WGS 84 datum, UTM Zone 37N), and pre-processed using the ERDAS IMAGINE 2014 software. The six major land uses were cropland, grassland, shrubland, forest, built-up land, and bodies of water, so we focused on these uses in our change-detection analysis. We used a hybrid land-use classification technique that combines unsupervised classification (using the ISODATA clustering algorithm) with supervised classification to assign pixels in the image to one of the land-use classes. We used field data gathered through field observations, interviews with key informants, and focus group discussions to validate the land-use classification and identify soil conservation measures. Four focus group discussions, each group comprised of 8 to 10 persons, and several indepth interviews with 18 key informants were held between November 2016 and May 2017. The group discussions and interviews were carried out by the corresponding author. Interviewed individuals included village administrators, watershed management committee members, and agricultural experts who had deep first-hand knowledge about the landuse types and their change history, watershed degradation and



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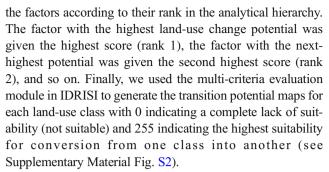
management practices, farming systems, and other socioeconomic and biophysical aspects of the watershed.

Land-use change modeling

We used the CA-Markov model implemented in the IDRISI Selva software (Eastman 2012) to simulate the future land-use patterns in 2050. We used Landsat-based land-use maps from 1986, 2001, and 2017 for model validation and creation of transition potential maps. In this context, transition potential maps represent the probability of a pixel for conversion into a particular land-use type. Land-use changes result from complex and strong interactions between people and their ecological and social environments (Lambin et al. 2001). Therefore, it is important to consider the set of potential factors and constraints during the creation of the transition potential maps. In this study, we selected biophysical parameters (e.g., elevation, slope, soil fertility) and accessibility factors (e.g., distances from roads, rivers, and towns) based on the knowledge and experience of the researchers, a review of the scientific literature, and the opinions of the farmers and local experts. For example, steep slopes limit farming practices and the expansion of built-up areas; however, such areas can be suitable for afforestation.

The remnant natural forests, existing water surfaces (e.g., the Rib Reservoir), and major road networks were considered to be constraints and were excluded from future land-use change. Fig. S1 and Table S1 in the Supplementary Material summarize the potential factors and constraints considered in our land-use modeling. We evaluated the contribution of the variables to predict land-use change by computing Cramer's V statistic, a measure of the association between variables, with a value ranging from 0 (no association) to +1 (complete association). We found moderate V values (0.25 to 0.58) the variables under consideration. We used the fuzzy set membership function approach provided by IDRISI to generate the individual suitability maps. Fuzzy membership standardizes continuous factors into the desired range using one of a number of membership function shapes (i.e., sigmoid, J-shaped, linear, user-defined) and types (i.e., monotonically increasing, monotonically decreasing, symmetrical) (see Supplementary Material Table S2). The fuzzy membership functions influence the probability of transition to a given land-use class by governing the shape and direction of the changes based on the values (control points) allocated to the factor images (Eastman 2012). For example, choosing a monotonically decreasing function would mean that the smallest distance value has the highest probability score, and as the distance increases, probability becomes smaller.

After defining the fuzzy set membership functions, we used the analytical hierarchy process (Saaty 1980) to consider the importance of the variables in land-use change (see supplementary Material Table S2). We applied the weight to



The accuracies of the simulated land uses were evaluated based on the kappa index and visual inspection. The overall kappa was 93.5%, which we considered to be satisfactory for the future land-use prediction.

Climate change modeling

Both the baseline and future climate data were analyzed in this study. The baseline climate data were used to estimate the baseline soil loss and validate the predicted climate scenarios. The daily precipitation and temperature records from 1973 to 2017 obtained from the Ethiopian National Meteorological Agency (www.ethiomet.gov.et) were used for baseline climate analysis. We considered five commonly used global circulation models (see Supplementary Material Table S3) using the moderate GHG concentration scenario (RCP 4.5) for future climate analysis. We evaluated the performance of the selected models by comparing the simulated rainfall data with observed data from 1986 to 2017 using three statistical measures: the root-mean-square error, coefficient of variation, and correlation coefficient (see Supplementary Material Table S4). Among the models evaluated, HadGEM2-ES performed the best in terms of capturing well-distributed rainfall and temperature values; we, therefore, selected it for use in the present study. The HadGEM2-ES model is a coupled Earth system model developed by the Met Office Hadley Centre for the Coupled Model Intercomparison Project Phase 5 (CMIP5) centennial simulations (https://portal.enes.org/). We downloaded the downscaled products of the HadGEM2-ES model for the period from 2018 to 2050 from the MarkSim data distribution center (http://gismap.ciat.cgiar.org/ MarkSimGCM/) for subsequent analysis.

Soil erosion modeling

We used the RUSLE model for erosion modeling. RUSLE is the revised form of the USLE model (Wischmeier and Smith 1978), with substantial improvements in its ability to account for the effects of soil surface roughness and local weather on the prediction of soil loss and sediment delivery. The RUSLE model can be integrated with the IDRISI software to enable estimation of soil loss for groups of pixels based on the slope, orientation, and slope length, which can be adjusted by the



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user (Chen et al. 2008). Such a slope-based estimation of soil loss is particularly useful in areas like the Rib watershed, where the slope strongly affects the development of rapid and erosive runoff. RUSLE requires six parameters, as follows:

$$A = R \times K \times LS \times C \times P \tag{1}$$

where A is the average soil loss per unit area (t ha⁻¹ year⁻¹), and the remaining factors are coefficients: R for runoff due to rain (MJ mm ha⁻¹ h⁻¹ year⁻¹), K for soil erodibility (t ha h MJ⁻¹ ha⁻¹ mm⁻¹), LS for topography, C for the soil and land cover, and P for conservation practices. The calculation methods we used for these coefficients are presented in the following sections.

Rainfall erosivity factor

R represents the erosive force of rainfall as a function of the rainfall volume, intensity, and duration (Alexakis et al. 2013). In the RUSLE model, the rainfall erosivity is estimated using the total storm energy and 30-min intensity (EI_{30}) of the rainfall (Renard et al. 1997). However, computation of the R values based on such intensive rainfall records is unrealistic in areas such as our study site that cannot provide data with this temporal resolution. Instead, we adopted an empirical formula developed by Hurni (1985) to estimate R values in Ethiopia:

$$R = -8.12 + (0.562 P) \tag{2}$$

where R is rainfall-runoff erosivity and P is the mean annual precipitation. To compute the R values, we converted the monthly rainfall values for the baseline period (1973 to 2017) and the future scenario (2018 to 2050) into the mean annual rainfall. We then developed iso-erosivity maps (Fig. 2a, b) by interpolating the mean annual rainfall at the meteorological stations shown in Fig. 1 to the entire watershed in the ArcGIS software using an ordinary kriging method based on a Gaussian function.

Soil erodibility factor

K represents the effect of soil properties and soil profile characteristics on soil loss. It can be best obtained from experimentally measured soil properties such as the soil's permeability, structure, and organic matter content. However, the absence of such detailed soil property data in most developing countries, including Ethiopia, often makes it infeasible to use empirical *K* values. As an alternative, researchers (e.g., Bewket and Teferi 2009; Brhane and Mekonen 2009) have suggested values of *K* for use in Ethiopia based on the soil color, which is believed to be a good proxy for soil physical properties. The *K* values in the present study were therefore

generated based on the soil color. The red, brown, and black soils were identified, and the corresponding K values were 0.25, 0.20, and 0.15, respectively (Fig. 2c).

Topographic factor

LS represents the effect of topography on soil loss, where L represents the effect of slope length on erosion and S reflects the influence of slope on erosion. Various empirical formulas have been developed around the world to determine the value of LS. In the present study, we generated an LS map of the study area based on a 30 m-resolution Shuttle Radar Topography Mission (SRTM) digital elevation model downloaded from the US Geological Survey data center. We used the equation of Moore and Burch (1986) to calculate the LS factor in ArcGIS (ESRI 2011).

$$LS = \left(F_{ac} \times \frac{CS}{22.13}\right)^{0.4} \times \left(\frac{\sin(\text{slope})}{0.0896}\right)^{1.3}$$
 (3)

where L is the slope length factor, S is the slope steepness factor, $F_{\rm ac}$ is a raster-based flow accumulation, and CS is the grid resolution. $F_{\rm ac}$ was computed from the digital elevation model using the Arc Hydro tools in ArcGIS (Fig. 2d).

Cover management factor

C represents the effects of cropping and management practices on soil erosion (Wischmeier and Smith 1978). Its value is largely controlled by the surface vegetation, land use, soil surface roughness, and soil moisture (Farhan and Nawaiseh 2015). With values ranging from 0 (well-protected soils) to 1 (bare soil), C is strongly related to land use (Ferreira et al. 2016). In this study, C values were determined based on the most recent (2017) and predicted future (2050) land-use data. We used the methods of Hurni (1985), Bewket and Teferi (2009), and Haregeweyn et al. (2017) to assign C values to each land-use class. The spatial distributions of C for the current and future scenarios are given in Fig. 2e–g.

Conservation practice factor

P represents the ratio of soil loss after the implementation of specific conservation practices to the corresponding soil loss before these practices (Renard et al. 1997). It is primarily used to assess the effects of farming practices and soil and water conservation measures on soil loss (Taye et al. 2018). The values of P range from 0 (well-conserved areas) to 1 (no conservation practice). Our field surveys indicated that soil and water conservation activities are not widely practiced in the Rib watershed. On the other hand, the farmers commonly use contour plowing to reduce erosion from surface runoff. In areas where contour plowing is widely practiced, Hurni



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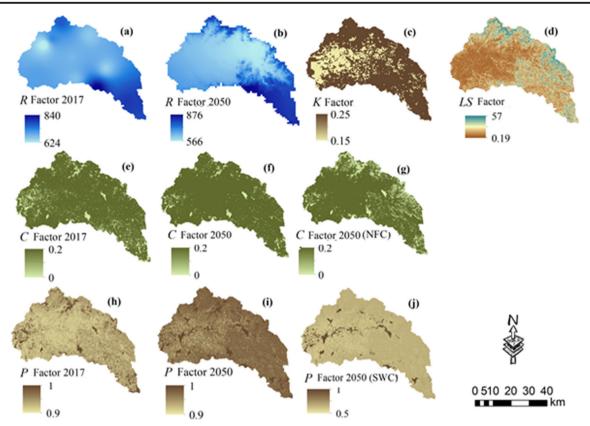


Fig. 2 Spatial distributions of the RUSLE input layers. **a** R factor for the baseline (1973 to 2017) condition, **b** R factor for the future scenario (2018 to 2050), **c** K factor, **d** LS factor, **e** C factor for the baseline condition, **f** C factor for the future scenario after

expansion of the native forest cover (NFC, scenario 3), \mathbf{h} P factor for the baseline condition, \mathbf{i} P factor for the future scenario, and \mathbf{j} P factor for the future scenario after the application of soil and water conservation measures (scenario 4)

(1985) and Taye et al. (2018) suggested using P values of 0.9 and 1.0 for agricultural and non-agricultural lands, respectively. We used a similar technique in the present study to assign P values to the whole watershed during the baseline period (2017). However, for the future scenario (2050), we chose areas with a slope greater than 20% as target areas for soil conservation practices. We, therefore, assigned a P value of 0.5 (Hurni 1985) to these areas. For areas with a smaller slope, we used the value from the baseline period. The spatial distributions of P in the current situation and future scenarios are given in Fig. 2h–j.

Scenarios and their descriptions

Based on the current and future climate and land-use datasets, we designed one baseline scenario (based on the conditions in 2017) and four future scenarios.

Scenario 1: baseline

This scenario assesses the current rate of soil loss in the watershed based on the current climate, land use, and management practices. The *R* values were computed from historical (1973 to 2017) rainfall data collected at gauging stations

located in and around the watershed, whereas the C and P values were generated based on the most recent (2017) landuse data.

Scenario 2: climate and land-use change

This scenario predicts future soil erosion under the influence of simultaneous climate and land-use changes under the current land-use regimes, without applying any conservation measures. We used mean annual rainfall (2018 to 2050) simulated by the HadGEM2-ES model to generate the map of R values, but generated the map of C values using the land use projected by the CA-Markov model for the year 2050.

Scenario 3: expansion of native forest cover

This scenario assesses the effectiveness of afforestation for reducing future soil erosion. Erosion rates in well-vegetated and protected areas are relatively low, ranging from 0.004 to 0.050 t ha⁻¹ year⁻¹ (Lal 1994). Forest reduces the rate of soil erosion by stabilizing soils and enhancing their water-holding capacity. Thus, the expansion of native forest cover should significantly reduce the future rate of soil loss in the study area. Based on our field surveys, the current forest type is



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dominated by eucalyptus species with little understory vegetation, which protects the soil poorly, leading to high erosion rates. In contrast, the native forest can have dense understory vegetation; as a result, soil erosion is typically low. However, much of the remaining native forest is seriously degraded because of grazing, and can no longer provide the same degree of protection offered by the intact forest.

In this scenario, we propose increasing the current native forest cover (<3% of the area) to 10.5% in 2050, mainly in hilly areas with a slope greater than 40%.

Scenario 4: application of physical soil conservation measures

This scenario evaluates the effectiveness of physical soil conservation measures in reducing soil loss. If effectively implemented, physical soil and water conservation measures such as sloping stone bunds, terraces, and check dams reduce future soil erosion in areas with steep slopes. These types of conservation measures have been widely practiced in northern highland Ethiopia (Tigray Region) and found to be effective in reducing surface runoff, facilitating water infiltration, and minimizing soil erosion (Gebrernichael et al. 2005; Nyssen et al. 2008; Taye et al. 2018). However, Moges and Taye (2017) reported that 10% or less of the total watershed currently includes soil and water conservation structures, and most of the existing structures have not been maintained properly. In this scenario, we assumed the implementation of physical soil conservation practices in all areas with a slope greater than 20% (which cover about 34% of the total catchment as the target zones for the application of physical soil conservation practices).

Scenario 5: combined scenarios

This scenario aimed to reduce the rate of soil loss by combining expanded native forest cover (scenario 3) with the use of soil and water conservation measures (scenario 4). We hypothesized that this combination would reduce soil erosion

Fig. 3 Framework for the five scenarios used to estimate soil erosion

better than using the two approaches separately. Figure 3 presents the overall methodological framework used for this study.

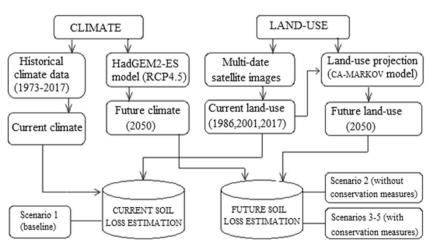
Results

Projected climate change

Compared with the baseline (1973 to 2017) scenario, the mean annual rainfall in the watershed is expected to increase by 17.3% by the year 2050 (Table 1). The model also predicted that the minimum and maximum temperatures would increase by 2.11 and 2.66 °C (25.5 and 12.2%), respectively. This agrees with previous studies (e.g., Conway and Schipper 2011; Kassie et al. 2014; Abera et al. 2018; Fentaw et al. 2018) that predicted increasing precipitation and temperatures in other parts of Ethiopia.

Land-use change

Our model predicts extensive land-use changes in the watershed between 2017 and 2050 (see Supplementary Material Table \$5). The built-up areas increased from 21 km² in 2017 to 56.5 km² in 2050 (a 169% increase). The cropland showed a relatively modest increase, from 1685 km² in 2001 to 1736.5 km² in 2017 (a 3% increase), and is projected to increase further, to 1772 km² in 2050 (a 2% increase). This relatively small growth can be explained by the fact that almost all potential agricultural areas have already been exploited, and the remaining areas are unsuitable for cultivation. The cropland expansion will take place only at the expense of grasslands and shrublands, which are projected to decrease in area by 66 and 25%, respectively, between 2017 and 2050. The forest cover showed a slight but continuous increase throughout the study period.





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Table 1 The mean annual rainfall and minimum and maximum temperatures in the current baseline situation and in the future scenario

Scenarios	Rainfall		Minimum temperature		Maximum temperature	
	Mean annual (mm)	Change (%)	Mean min. (°C)	Change (%)	Mean max. (°C)	Change (%)
Baseline	1351	_	8.27	_	21.78	
Future	1585	17.3	10.38	25.5	24.43	12.2

Modeled soil erosion

Table 2 presents the mean and total soil loss rates corresponding to the five scenarios. The total annual soil loss for the whole watershed ranged from 7.93×10^6 t year⁻¹ in 2017 to 9.75×10^6 t year⁻¹ in 2050 in scenario 2 (an increase of about 23%). The mean annual soil loss ranged from 69.6 t ha⁻¹ year⁻¹ in 2017 to 78 t ha⁻¹ year⁻¹ in scenario 2. The mean annual soil loss of 78 t ha⁻¹ year⁻¹ in 2050 (scenario 2) represents an increase of 12% compared with the baseline period (2017). However, the high mean annual soil loss predicted under scenario 2 could be reduced to less than 45 t ha⁻¹ year⁻¹ (a 42% decrease) by implementation of either conservation measure, with the greatest effect achieved by a combination of both measures in scenario 5 (soil loss reduced to 30.4 t ha⁻¹ year⁻¹, a decrease of 61%,).

To identify conservation priorities based on the erosion risk, we classified the estimated soil loss into six levels (FAO 1986; Molla and Sisheber 2017): slight, moderate, high, very high, severe, and very severe (Fig. 4). For all of the designed scenarios, the highest rates of soil loss were observed in the northern, eastern, and southeastern parts of the watershed, whereas the western parts and most of the central parts experienced lower soil loss.

Discussion

The effect of climate change on soil erosion

The mean annual soil losses estimated in this study agree with experiment-based estimates by Bewket and Teferi (2009), Yihenew and Yihenew (2013), and Zerihun et al. (2018), which ranged from 49 to 93 t ha⁻¹ year⁻¹. The variability in the spatial distribution of soil loss can be attributed to variability in the biophysical components of the watershed, including the topography, rainfall, land use (i.e., surface cover), and soil

type. The magnitude of soil loss in the watershed can be evaluated by comparing the estimated soil losses with threshold values that define the soil loss tolerance level. In all five scenarios, more than 50% of the watershed experienced erosion rates higher than the maximum tolerable level of 18 t ha⁻¹ year⁻¹ for Ethiopia (Hurni 1985).

Climate change influences soil erosion in different ways. The most direct and significant influence of climate change is through changes in the erosive power of rainfall (Nearing et al. 2004). Higher rainfall amounts and intensities (i.e., amount per unit time), as well as extreme rainfall events, increase soil loss due to their direct impacts on runoff rates and the mechanics of soil erosion (Li and Fang 2016). Pruski and Nearing (2002) predicted that every 1% change in rainfall amount could change runoff by 2% and erosion by 1.7%. Similarly, Zhang (2007) showed that a 4 to 18% increase in precipitation can increase runoff by 49 to 112% and soil erosion by 31 to 167%. Therefore, the predicted 17% increase in rainfall in the Rib watershed is likely to cause a proportionally high increase in runoff and erosion, especially when combined with other factors such as the watershed's poor vegetation cover and limited use of soil and water conservation measures.

Another potential effect of climate change on soil erosion relates to rising temperatures and atmospheric CO₂ concentrations. The predicted temperature increase can have both positive and negative effects on soil erosion. As temperature increases, evapotranspiration increases and soil moisture decreases, thereby increasing the soil's water infiltration capacity and reducing runoff and soil erosion (Xu 2003). Moreover, CO₂-driven changes in plant biomass production and organic matter decomposition rates may decrease soil erosion by increasing vegetation cover (Mullan 2013). In contrast, plant stress created by high temperatures may increase erosion rates by increasing evaporation, reducing the availability of soil water, and thereby decreasing plant growth (Pruski and Nearing 2002; Li and Fang 2016). An increase in temperature also increases organic matter decomposition rates due to

Table 2 The mean and total soil losses for current baseline scenario (scenario 1) and four future scenarios: 2, current trends without soil conservation measures; 3, expansion of native forest cover; 4, use of physical conservation measures; 5, combination of scenarios 3 and 4

Soil loss	Scenario 1	Scenario 2	Scenario 3	Scenario 4	Scenario 5
Mean (t ha ⁻¹ year ⁻¹)	69.6	78.00	44.50	38.00	30.4
Total (Mt year ⁻¹)	7.93	9.75	7.55	5.14	3.72



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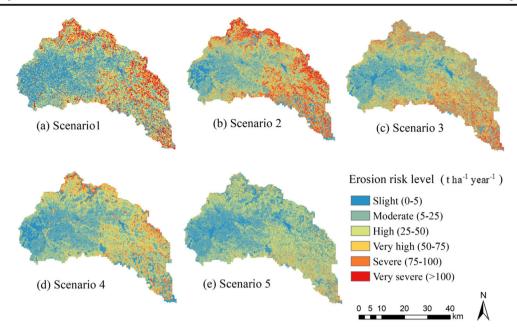


Fig. 4 Soil erosion risk maps corresponding to the five studied scenarios. **a** Current soil loss for the baseline conditions in 2017 (scenario 1), **b** future soil loss under the predicted climate and land-use changes but without any conservation actions (scenario 2), **c** future soil loss after expansion of native forest cover to reduce erosion (scenario 3), **d** future

soil loss after application of soil and water conservation measures (scenario 4), and ${\bf e}$ future soil loss under a combination of the conservation actions in scenarios 3 and 4 (scenario 5). Future values are predicted for 2050

increased soil microbial activity (Nearing et al. 2004), and this can expose the soil to surface runoff, leading to soil erosion. Williams et al. (1996) note that warm temperatures can also cause soil erosion through changes in land use and crop management practices to accommodate the new climatic regime.

Overall, our simulations suggest that in the absence of measures to reduce runoff, soil erosion in the study area is expected to become more severe, especially in upland areas where soil erosion is already severe, in response to the direct influence of changes in rainfall patterns and the indirect influence of rising temperatures. Under the influence of the region's rapidly growing human population and limited access to advanced agricultural technologies, any climate-induced increase in future soil erosion would significantly decrease the productivity of rainfed subsistence agriculture and jeopardize the livelihood of the watershed's rural community.

The effect of land-use change on soil erosion

Land-use change is among the major factors that determine the rate of soil erosion. Rapid population growth, a shortage of arable land, and increased demand for food production have significantly intensified the pressure on the land and aggravated land-use change in most parts of highland Ethiopia, including the Rib watershed (Zeleke and Hurni 2001; Meire et al. 2013; Meshesha et al. 2014; Moges and Bhat 2018). Conversion of natural vegetation (native forest, shrubs, and grasses) to cropland was the major land-use change observed in the Rib watershed. Zhou et al. (2008) found that watersheds

tend to experience high rates of erosion when more than 30% of the soil is exposed, whereas vegetation cover of more than 78% can greatly reduce erosion by water. In the Rib watershed, however, less than 10% of the total area is covered by natural vegetation, even though this vegetation is located primarily in areas dominated by rugged topography and steep slopes, which are unsuitable for agriculture and construction of built-up areas.

Natural vegetation cover affects the processes that govern the rate of soil erosion in various ways. For example, this vegetation tends to do a better job of protecting the soil because the higher species diversity results in higher vegetation cover, greater deposition of organic matter (e.g., leaf litter), and more complete exploitation of the whole soil volume by plant roots than would be the case with artificial vegetation communities. Removing natural vegetation increases soil erosion by reducing the protection of the soil, litter production, and organic matter accumulation, leading to a higher value of C (Wijitkosum 2012). Litter and organic matter not only directly protect the surface soil from the kinetic energy of falling raindrops and the resulting splash erosion, but also conserve surface water (Wei et al. 2007). Converting natural vegetation to other land-use types also diminishes the role of plant root systems, which reduce soil erodibility by improving intrinsic properties of the soil such as aggregate stability, infiltration capacity, bulk density, texture, organic matter and chemical contents, and shear strength (Gyssels et al. 2005).

The rate of soil loss was highest in croplands followed by non-native forest and shrubland. The high soil loss in cropland



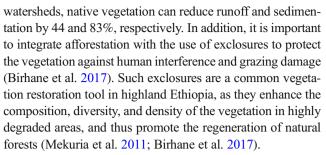
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results from widespread agricultural activities combined with poor soil conservation practices. Our field observations revealed that most of the croplands in the watershed are located in areas dominated by rugged terrain, leading to the cultivation of steep slopes; the exposed soil surface at such sites would greatly increase soil erosion compared with vegetated surfaces. The high erosion rates observed in the forested land have several possible explanations. In contrast with the findings of studies in other parts of Ethiopia (e.g., Haregeweyn et al. 2012; Hassen and Assen 2018), we found a slight but continuous growth of forest cover during the study period. However, our field observations and interviews revealed that the increase in forest cover that we observed was not because of the growth of native forest cover, but rather was due to expansion of the area of eucalyptus plantations. Eucalyptus has been introduced in most parts of highland Ethiopia as an alternative to native species (Nyssen et al. 2009). Moges and Bhat (2018) report that these trees are the main source of income and fuelwood in the Rib watershed, which strongly encourages residents to introduce or expand the area of eucalyptus trees in their farmland. Soils under eucalyptus trees become more acidic and have lower organic matter and nutrient contents than the soils under native forests (Liang et al. 2016), leading to low soil porosity, high surface runoff, and severe erosion. Therefore, it is essential to convert eucalyptus plantations to native forests to improve soil conservation in the watershed.

Another possible reason for the high soil loss in forests is that the few remnant native forests in the watershed are located in hilly areas with steep slopes. This topography considerably increases soil erosion due to its direct influences on surface runoff, drainage, and sedimentation rates (Dessalegn et al. 2014). Haigh et al. (1995) suggest that a minimum of 60% of forest cover is required to effectively prevent soil erosion in forested land. In the Rib watershed, however, the residual native forests are sparsely vegetated and mostly used for cattle grazing, which can also exacerbate erosion due to the removal of vegetation by grazing and the damage to the soil structure caused by the hooves of the cattle.

The effectiveness of potential conservation measures

To test the effectiveness of the soil conservation measures, we designed two conservation scenarios and evaluated their effectiveness, alone or when combined, for reducing future soil erosion. We found that expansion of the native forest cover (scenario 3) can reduce the total annual soil loss by more than 22% compared with future soil loss rates without any soil conservation measures (scenario 2). This agrees with the findings of Tesfahunegn et al. (2012), who demonstrated that afforestation was capable of reducing runoff and sediment yields by up to 51% in highly degraded areas of northern Ethiopia. Zhang et al. (2016) reported that in small



Our results also show that the construction of physical soil and water conservation measures such as terraces and sloping stone bunds (scenario 4) can reduce total soil loss of the watershed by 47% compared with the soil loss with no soil conservation measures (scenario 2). Such physical measures not only intercept runoff, increase infiltration, and trap sediments, but also reduce the slope of the land between adjacent structures, thereby slowing runoff and reducing its erosive power (Nyssen et al. 2007). An experimental study by Melaku et al. (2018) indicated that soil and stone bunds can notably reduce the slope length where they are implemented within a watershed and provide retention space for sediments; as a result, they gradually reduce the original slope and make it more suitable for cultivation by trapping nutrient-rich-eroded sediments. Terraces work similarly. Combining terraces and stone bunds with suitable crops can further reduce the ability of a steep slope to increase soil erosion (Descheemaeker et al. 2006). In addition, Lemessa (2001) and Woldemariam et al. (2018) reported that khat (Catha edulis), a perennial cash crop, can potentially reduce soil erosion by minimizing the rainfall's runoff velocity in the highlands of Ethiopia.

The total annual soil loss estimated under scenario 2 could be reduced by 62% by combining afforestation with the use of physical soil and water conservation measures (scenario 5). On this basis, we recommend that local governments encourage the combined use of the two conservation measures to reduce future soil loss rather than applying them separately. Tesfahunegn et al. (2012) found that the application of soil conservation measures alone was less effective than combining these measures with increased afforestation. Application of soil conservation measures such as intercropping, stone bunds, and afforestation individually is no longer sufficient to sustain landscape management efforts, but combining several of these nature-based solutions can potentially enhance the soil's health and its functions, thereby maintaining or restoring the provision of ecosystem services (Keesstra et al. 2018). It is important to note that although we found some evidence of soil conservation measures being used in our field research, the physical structures were often poorly maintained. Thus, to ensure that these structures provide their full potential benefit, local governments or the national government should increase agricultural extension activities to teach farmers the importance of maintaining these structures and should consider subsidizing the cost of this maintenance.



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Limitations of our soil erosion model

GIS tools are useful for modeling complex relationships between land use, soil erosion, and topography, and when they are combined with a soil loss model such as RUSLE, they can contribute greatly to sustainable land-use planning (Kull et al. 2005). However, several factors create uncertainty in our modeling results, including the absence of reliable data to parameterize the models and uncertainties in the climate model outputs that drive the soil loss model (Alewell et al. 2019). The socioeconomic and political conditions in a study area will also strongly affect the land-use changes. Because we could not obtain such data for our study area, the predicted future land-use patterns were based on the biophysical characteristics of the watershed. Although climate models are becoming more accurate, predictions for small study areas differ between models (O'Neal et al. 2005). To reduce such discrepancies, it may be worthwhile to combine the outputs of multiple models instead of depending on a single model's outputs. In this study, however, we performed our climate predictions using a single model (HadGEM2-ES).

Another possible limitation of our study was the lack of measured sediment yields or soil loss rates to validate the simulated outputs. However, we accounted for the results of previous studies and for information gathered during our field observations to compare the severity and spatial distribution of the estimated and actual soil loss in the watershed. Meshesha et al. (2012) found good reliability and validity of RUSLE outputs, with an overall accuracy of erosion estimates of 71.1%. Further field-based studies of sediment transport and deposition could be used to calibrate RUSLE and improve the land use and management decisions in the study area.

Conclusions

In this study, we analyzed the potential impacts of climate and land-use changes on soil erosion and evaluated the potential effectiveness of simple conservation measures. Our modeling predicted considerable changes in climate and land use in the watershed by 2050. We predict a continuous and substantial increase in cropland at the expense of native forest, grassland, and shrubland. Precipitation and temperatures will increase by 2050, leading to direct and indirect effects that will increase soil erosion if preventive measures are not implemented. The rapid human population growth and high pressure that will be placed on limited watershed resources may increase future soil loss compared with the predicted rates. Therefore, it is essential to begin implementing conservation measures to minimize soil erosion and stabilize the landscape. Our modeling suggests that expansion of the native forest cover and implementation of physical soil conservation structures can greatly reduce soil erosion, but the effects will be greatest when these measures are combined. However, additional measures may be necessary to further reduce erosion.

Our results demonstrate that the combination of models we employed has a high potential for predicting future soil erosion and evaluating potential conservation measures to mitigate the effects of climate and land-use change. Our results can be used to increase the awareness of the farmers about the causes, magnitudes, and effects of soil loss and motivate them to investigate soil conservation measures. However, government investments and support may be required to encourage farmers to change their practices. Our results will also help decision-makers to improve their planning of future land use and to identify priority soil conservation to improve the sustainability of watershed management. Additional studies on the economic feasibility of the proposed conservation measures could help the farmers and decision-makers to better prioritize the optimal soil conservation areas and methods.

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